Unexpected enrichment of nitrogen during the diagenetic evolution of sedimentary organic matter

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This study examines the behaviour of nitrogen during coalification in two coal series: the sedimentologically homogeneous coal series of the Mahakam delta, Indonesia and a reference coal series composed of samples from several sources. The variations in the total nitrogen contents, in hydrolysable nitrogen, and in the natural isotopic abundance of ¹⁵N are reported. The behaviour of nitrogen during coalification presents three main stages: 1. a first stage where microbial and/or chemical degradations produce a loss of nitrogen; 2. a second stage of preservation, where organic nitrogen does not take part in the reactions involved and consequently remains unchanged in the residual organic matter; and 3. a third stage of thermal alteration, where nitrogen is mainly released with methane.

(Keywords: coal; coalification; nitrogen)

Compared with the data available on the character and distribution of carbon, hydrogen and oxygen in coals and kerogens, there is relatively little information on nitrogen, sulphur and phosphorus groups. At the beginning of coalification nitrogen occurs in humic like compounds and it might be thought that nitrogen groups, like oxygen groups, would disappear continuously during diagenesis. In fact, this Paper shows that nitrogen is preserved in sedimentary organic matter over a large part of the period of diagenesis.

A geochemical study of nitrogen in the sedimentologically homogeneous coal series of the Mahakam delta (Indonesia) is reported, and geochemical changes of nitrogen in this coal series are compared with those of another coal series. Data are related to the coal organic carbon, taken as a rank parameter which allows a comparison of the different evolution stages according to Stach et al.². The implications of variations in total nitrogen content for prediction of losses of organic matter and hydrocarbons during diagenesis are discussed.

The Mahakam delta is situated in Indonesia, on the eastern part of the Kutèi Basin, in south-eastern Borneo (Kalimantan). The geological setting and physical properties of Mahakam coals are extensively described elsewhere^{3–8}. This coal series is relatively uniform in terms of macerals as well as sedimentology: sedimentological conditions, climatic environment and the origin of the coal organic matter have always been constant during the delta history. Huminite or vitrinite represents 71–88% of these coals.

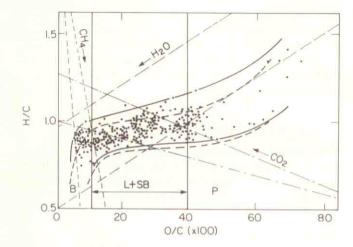
Mahakam samples, which range from lignite to bitu-

minous coal, were provided by the Socièté Total-Indonésie and the Compagnie Française des Pètroles. They were obtained from cuttings taken at various depth intervals from many boreholes to a maximum depth of 3500 m. As the number of samples was exceptionally high (600), sampling can be considered as representative over the coal series. Reference coals were provided by the Institut Français du Pétrole and by Dr H. W. Hageman and Dr M. Teichmüller. This reference set of samples does not represent all possible coal types (for instance, it contained no Gondwana coal samples).

Freeze-dried and crushed coals were analysed according to the methods of Durand and Monin⁹. Hydrolysis was performed in 6N HCl for 12 h, under reflux. Hydrolysable nitrogen was analysed according to the Kjeldhal method¹⁰. ¹⁵N was determined according to the method of Mariotti et al.^{11,12}.

Figure 1 shows the atomic H/C versus O/C diagram¹³ for the Mahakam coal series, and it is compared with the atomic H/C and O/C ratios of the reference collection of vitrinite-rich coals from several origins. This diagram offers an insight into the course of the process occurring during coalification. It shows that the coalification processes of the two coal series (Mahakam and reference) did not differ significantly.

In Figure 2 it is seen that amounts of total nitrogen in Mahakam coals and in the reference coals increase significantly during the lignite and the subbituminous stages, then start to decrease during the bituminous stage. In the reference coals, the nitrogen values quickly decrease during the anthracite stage. Scatter due to initial type of plant matter² is relatively low for the Mahakam coal series. The uniformity of the Mahakam coal series permits for the first time observation of an unexpected increase in



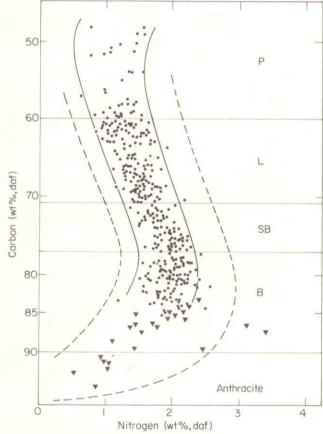


Figure 2 Nitrogen *versus* rank. Data are plotted as a function of coal organic carbon taken as rank parameter. Organic carbon enables us to set the P, L, SB, B and anthracite stages according to Stach *et al.*² The envelope (——) of the Mahakam samples (●) is narrower than the envelope (---) of a reference coal series, which does not extend into the lowest ranks. Coals of high rank studied by Schapiro and Gray¹⁵ (▼) follow the same downward trend as the Mahakam coals, as predicted at the bituminous stage

nitrogen percentages over the lignite and subbituminous stages, which up to now was less clear in other published work, because of the scatter within the data.

An explanation of this increase may simply be that the proportion of nitrogen increases because of the loss of more labile oxygen. This hypothesis is in accordance with the following observations. Hydrolysable nitrogen drops sharply at 51 wt% C (daf), i.e. well down in the peat stage. At the beginning of the lignite stage, total nitrogen is no longer hydrolysable, as seen in *Table 1*. Very unstable compounds would be completely decomposed and would furnish ammonia during early diagenesis. This decomposition could account for the early decrease in nitrogen percentages shown in *Figure 2*. The main part of the organic nitrogen is stable in lignites, and must be referred to as heterocyclic¹.

Figure 3 shows a larger scattering of $\delta^{15}N$ values for lower rank Mahakam coals, compared with the higher ranks, where $\delta^{15}N$ values are not so dispersed. Given the homogeneity of the Mahakam coals, it may be that the shift in $\delta^{15}N$ observed during coal maturation are significant. If nitrogen loss, such as is expected during the peat and bituminous stages, is able to produce an isotopic fractionation, a partial proof of the preservation of the nitrogen of coal during the lignite and subbituminous stages may be inferred from the invariability of the $\delta^{15}N$ values during these stages.

Increase of nitrogen percentages is important for an understanding of organic matter diagenesis, as well as for establishing more clearly the mode of nitrogen evolution in coals of increasing rank. If it is assumed that nitrogen is accumulated in coal, the total loss of organic matter can be calculated and a quantitative model of evolution can be computed 14. If the nitrogen percentages at the beginning and end of a stage are N1 and N2, respectively, weight loss percentage is equal to 100(1 - N1/N2); this averages 24% during the lignite stage, and 7% during the subbituminous stage. Since N1 and N2 will not be known with a great degree of accuracy ($\pm 5\%$), the weight loss is not known very precisely, but at least an estimate can be made of the orders of magnitude and the amounts of products which are formed during diagenesis can be calculated.

Release of nitrogen occurs at the beginning of the bituminous stage, but is particularly noticeable at the anthracite stage, as shown for an anthracite series

Table 1 Hydrolysable nitrogen contents of representative samples of Mahakam coal series. Each representative sample is a mixture of several individual samples (fresh vegetal samples, litter samples, buried samples having a same huminite-vitrinite reflectance)

Total organic carbon in coal (wt %, daf)	Hydrolysable nitrogen (mg/100 mg nitrogen)
48.8	55.1
49.0	53.8
50.1	51.5
51.6	16.5
56.6	6.4
57.7	3.5
57.7	3.9
59.4	4.3
60.8	2.7
61.1	3.8
61.6	3.3
62.5	3.2
66.7	1.9
68.3	1,5
69.4	1.1

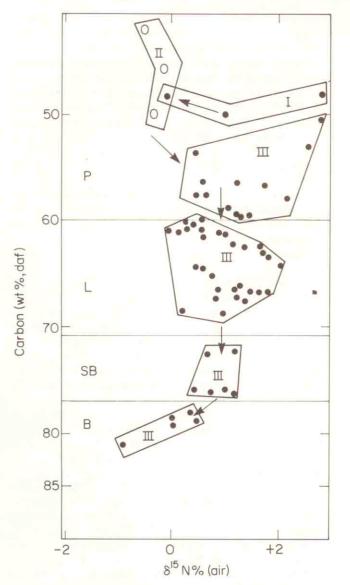


Figure 3 δ^{15} N for Mahakam coals. Data are plotted *versus* organic carbon percentages, as in *Figure 2*. (I) refers to the area of fresh vegetation (*Nypa fruticans*, a characteristic species of the deltaic vegetation), (II) refers to the area of litter (detritus of various plants from the surface deposit), (III) refers to areas of buried samples (cuttings obtained at various depth intervals)

(Figure 2)¹⁵. The mechanism involved in the process of thermal nitrogen elimination could also be responsible for gas formation¹⁶, which begins later in the geothermal sequence than does oil formation. This would explain the presence of nitrogen in some gas inclusions of metamorphized rocks²¹, and in some gas accumulations¹⁷⁻²⁰.

This research has been commenced because petroleum geologists are just beginning to drill into the major deposits of gas below the level at which oil forms. Further studies are now in progress to establish the mechanism of nitrogen loss and its connection with gas accumulation.

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